



CAMAERA

GA 2025, Lille

Parametrizations of deposition velocity for particles

Rostislav Kouznetsov, Mikhail Sofiev, Andreas Uppstu, Risto Hänninen
Finnish Meteorological Institute

CAMAERA GA 2025 4.6.2025



PROGRAMME OF
THE EUROPEAN UNION



IMPLEMENTED BY



1



Coordinated by

HYGEOS

Deposition velocity: model parameter

$$\text{Flux} = V_d(\text{particle, flow, surface}) \cdot \text{Concentration}$$

- V_d implies that flux is proportional to concentration
- V_d to be parameterized (experiments, mechanistic models etc.)
- Often u_* used to get velocity dimension for V_d

Direct method

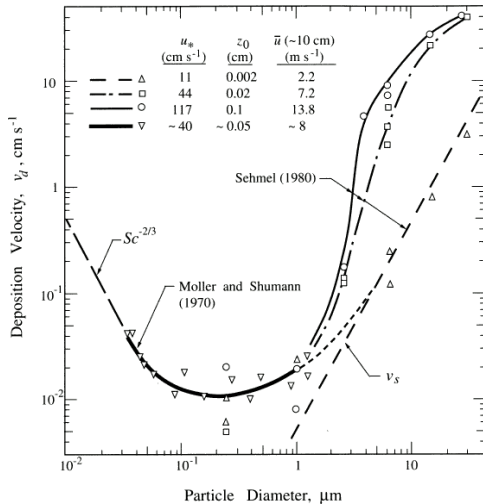
- Observe concentration and deposited amount
- Require both measurements
- Laborious

Indirect method

- Observe concentration and in-air vertical flux
- Any flux measurements
- Assumption: vertical flux is deposition flux



Slinn, 1980



Slinn (1978), Fig. 9, after SP98

As in Seinfeld & Pandis, 1998

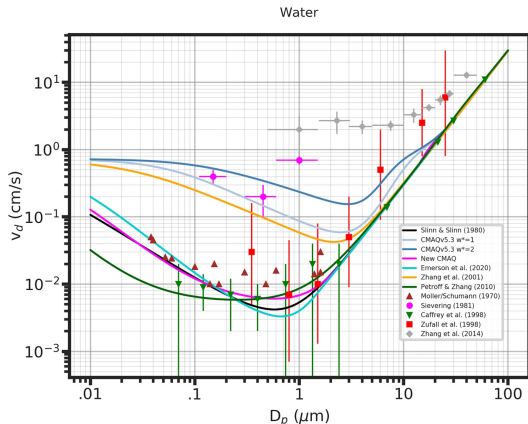
- Suggested by [Slinn&Slinn \(1980\)](#) for water surfaces
- Concepts of r_a , r_b and virtual resistance

$$v_d = \frac{1}{r_a + r_b + v_s r_a r_b} + v_s$$

- r_b accounts for diffusion and impaction
- No surface-specific parameters
- Very easy to implement
- Recommended by SP1998 for all surfaces
- Criticized by [Venkatram and Pleim \(1999\)](#)
- Much lower V_d than measured over nature

Zhang et al, 2001

- Resistance analogy
- Surface-dependent r_b
- Add interception term (Slinn, 1982)
- 4 parameters (3 deposition-specific)
- 15 LUC, 5 seasons
- x10 higher V_d than Slinn (1980)
- Recommended by SP2006, and SP2016



Pleim et al (2022), Fig. 6

LUC		1	2	3	4	5	6	7	8	9	10	11	12	13	14	15
Z (m)	SC 1	0.8	2.65	0.85	1.05	1.15	0.1	0.1	0.04	0.03	0.1	0.03	0.01	$f(u)$	$f(u)$	1.0
	SC 2	0.9	2.65	0.85	1.05	1.15	0.1	0.1	0.04	0.03	0.1	0.03	0.01	$f(u)$	$f(u)$	1.0
	SC 3	0.9	2.65	0.80	0.95	1.15	0.05	0.02	0.04	0.03	0.1	0.02	0.01	$f(u)$	$f(u)$	1.0
	SC 4	0.9	2.65	0.55	0.55	1.15	0.02	0.02	0.04	0.03	0.1	0.02	0.01	$f(u)$	$f(u)$	1.0
	SC 5	0.8	2.65	0.60	0.75	1.15	0.05	0.05	0.04	0.03	0.1	0.03	0.01	$f(u)$	$f(u)$	1.0
A (mm)	SC 1	2.0	5.0	2.0	5.0	5.0	2.0	2.0	na	na	10.0	10.0	na	na	na	10.0
	SC 2	2.0	5.0	2.0	5.0	5.0	2.0	2.0	na	na	10.0	10.0	na	na	na	10.0
	SC 3	2.0	5.0	5.0	10.0	5.0	5.0	5.0	na	na	10.0	10.0	na	na	na	10.0
	SC 4	2.0	5.0	5.0	10.0	5.0	5.0	5.0	na	na	10.0	10.0	na	na	na	10.0
	SC 5	2.0	5.0	2.0	5.0	5.0	2.0	2.0	na	na	10.0	10.0	na	na	na	10.0
α		1.0	0.6	1.1	0.8	0.8	1.2	1.2	50.0	50.0	1.3	2.0	50.0	100.0	100.0	1.5
γ		0.56	0.58	0.56	0.56	0.56	0.54	0.54	0.54	0.54	0.54	0.54	0.54	0.50	0.50	0.56

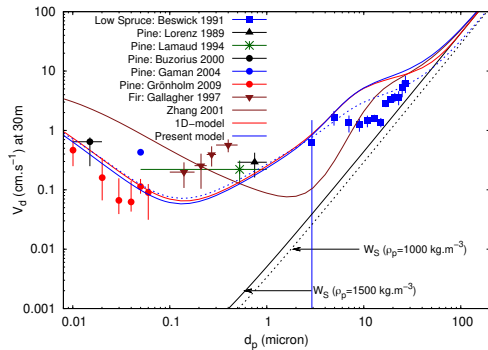
Note: $f(u)$ represents a function of wind speed (u) and na represents not applicable.

Zhang et al (2001), Table 3

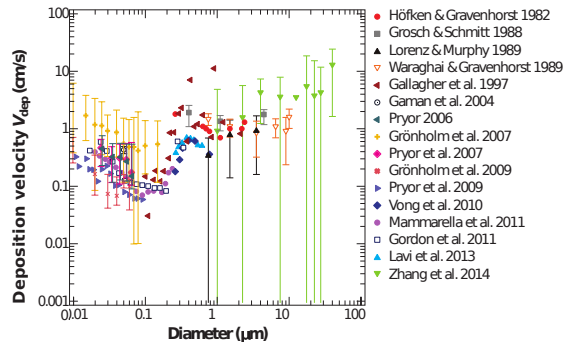
Petroff and Zhang, 2010

- “Exponential” form for aerodynamic layer
- Separate smooth surfaces, only diffusivity and settling
- 10 parameters (6 deposition-specific), 26 Land-Use Categories
- Better fits measurements

Deposition velocity on coniferous forests



Petroff and Zhang, 2010, Fig. 5

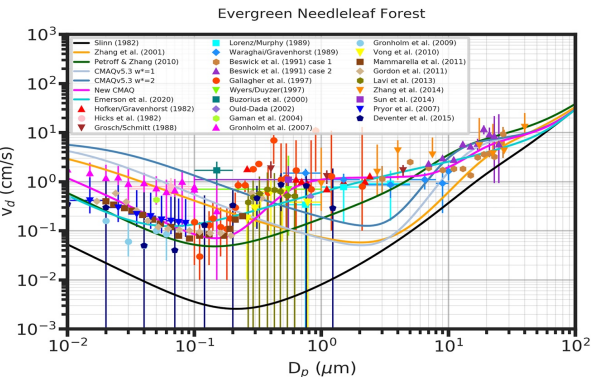


Farmer et al., 2021, Fig. 5

Recent schemes

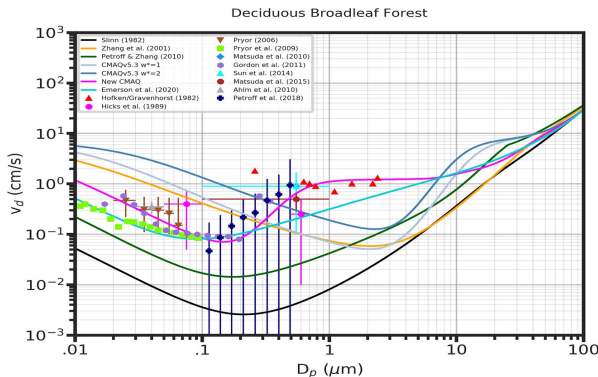
Emerson et al., 2020

- Driven by observations
- Zhang 2001 with adjusted parameters
- Interception $\sim (d_p/A)^{0.8}$ (Sic!)
- Shifts deposition minimum to $0.1\mu\text{m}$
- Cyan line



Pleim, 2022

- Two collector scales for impaction
- Impaction on leaf hairs (Slinn, 1982)
- Interception $\sim (d_p/A)^2$, but negligible
- Sharper deposition minimum $0.1\mu\text{m}$
- Magenta line

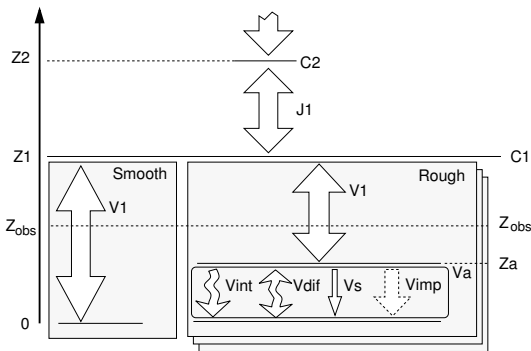


KS2012 scheme

- “Exponential” scheme for finite layers
- Pericles with D , d_p , τ rather than ‘size’
- Asymptotic transition to gases
- Smooth/rough: roughness Reynolds number

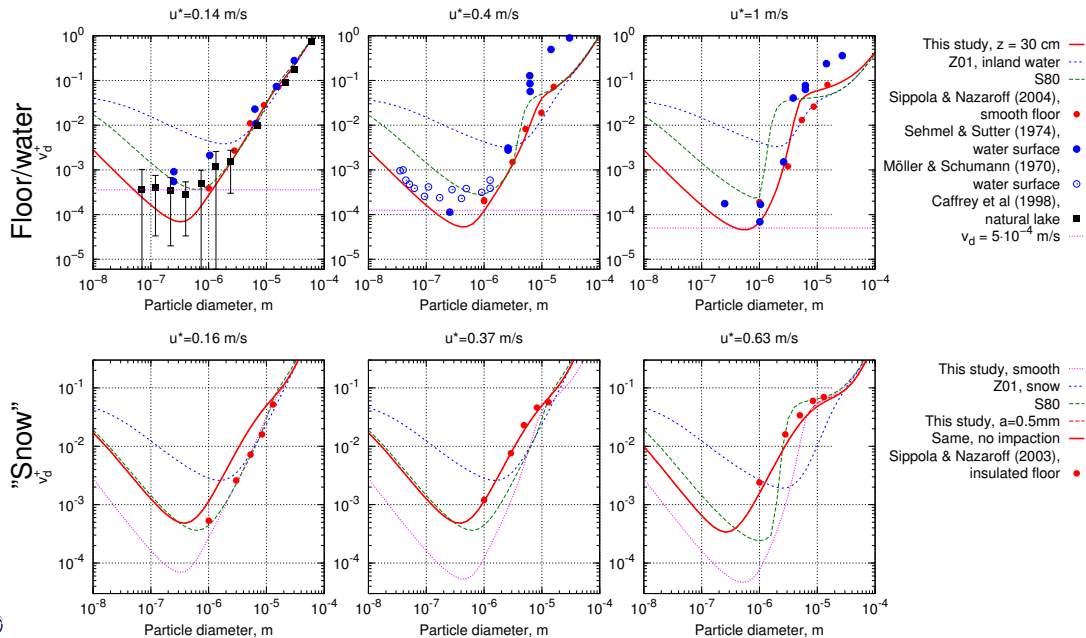
$$u_* z_0 / \nu \sim 3$$

- Smooth surfaces
 - Rigorously derived scheme
 - Settling, turbophoresis, diffusion
 - No fitting parameters
- Rough surfaces
 - Filtration in canopy top
 - Particle vs momentum efficiency
 - Settling, impaction, interception, diffusion
 - One deposition-specific scale a
 - Can be calibrated with ‘collectors’ rather than with ‘canopy’
- Phoretic processes (thermophoresis so far)

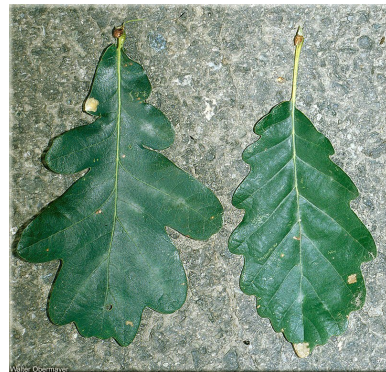
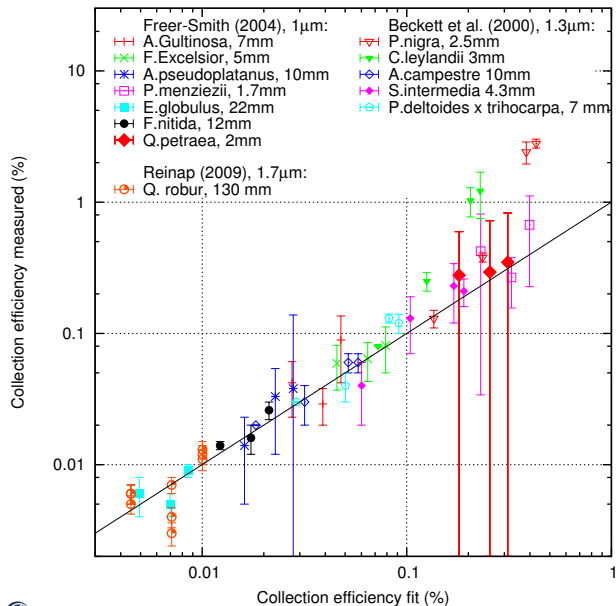


Kouznetsov & Sofiev 2012, JGR

Smooth surfaces and dynamic switch



Collector size from WT: trees/branches (slide from ITM 2013)



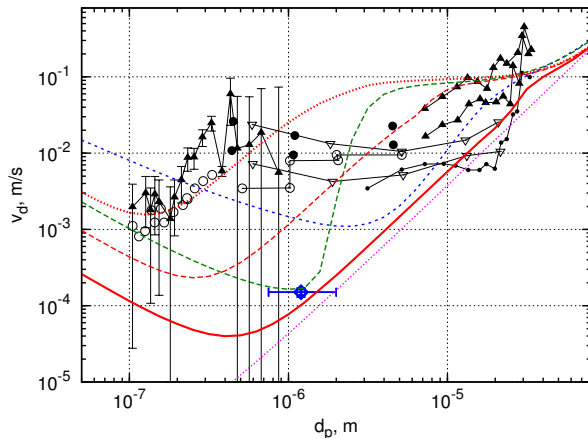
Courtesy of Walter Obermayer

Quercus Robur & Quercus petraea

Problems:

- Few studies
- Wide and uncertain particle size spectra

High vegetation (slide from ITM 2013)



- | | |
|----------------------------------|----------------------------------|
| ○ Joutsenoja (1992) | ◆ Reinap & Wiman (2009) + v_s |
| ⊢ Gallagher et al. (1997) | --- Z01, broadleaf forest |
| ▲ Gallagher et al. (1992) | --- S80 |
| ● Grosch & Schmidt (1992) | — This study, $a=7\text{mm}$ |
| ⊢ Bewsick et al. (1991) | --- This study, $a=0.5\text{mm}$ |
| ⊖ Lorenz & Murphy (1989) | ⋯ This study, $a=0.05\text{mm}$ |
| ▽ Waranghai & Gravenhorst (1989) | ⋯ Settling, v_s |

Probable causes of discrepancies:

- Electricity?
 - Boltzmann charging: no sharp size dependence
 - Dipole charging: 10 orders of magnitude weaker
- Wrong measured size?
- Complex particle shapes?
- Humidity?
- Something else?

Black and white symbols Gallagher (1997)

Deposition on water

V_d for $1\text{ }\mu\text{m}$ particles

- Direct measurements: $V_d \sim 0.1\text{ mm/s}$
- Mechanistic models: Same
- Indirect measurements:
 $V_d \sim 10\text{ mm/s}$
- 2 orders of magnitude difference!

Experiments and parametrisations

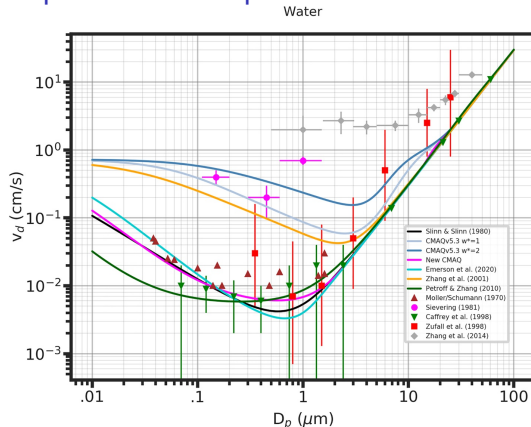


Fig. 6 from Pleim et al (2022)

Deposition on grass

V_d for $1\text{ }\mu\text{m}$ particles

- Direct measurements: $V_d \sim 0.1\text{ mm/s}$
- Mechanistic models: Same
- Indirect measurements:
 $V_d \sim 10\text{ mm/s}$
- Again orders of magnitude difference
- V_d is one of many uncertainties in transport models

Experiments and parametrisations

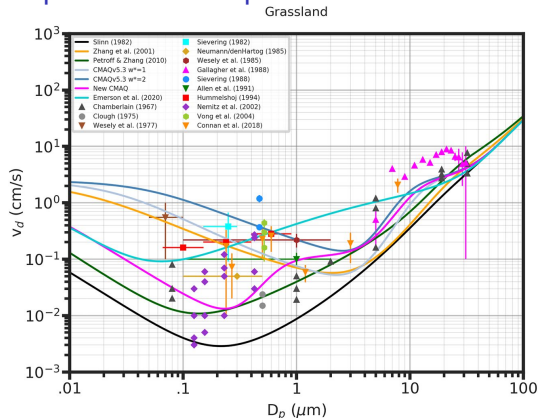


Fig. 4 from [Pleim et al \(2022\)](#)

Deposition on needle-leaf forest

V_d for $1\text{ }\mu\text{m}$ particles

- Measurements: $V_d \sim 10\text{ mm/s}$
- Impressive consensus
- Mechanistic models: $V_d \sim 0.1\text{ mm/s}$
- "Ground truth" often preferred
- Is it a clear case?
- All observations are indirect

Can we simulate high ' V_d ' with KS2012?

Experiments and parametrisations

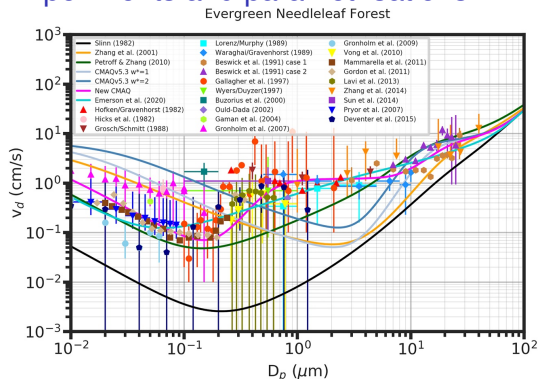


Fig. 2 from Pleim et al (2022)

Fluxes with mechanistic model?

Measurements (Gallagher, 1997)

- Fluxes ($0.1 - 0.4 \mu\text{m}$) over forest canopy
- Eddy Covariance with OPC

SILAM simulation

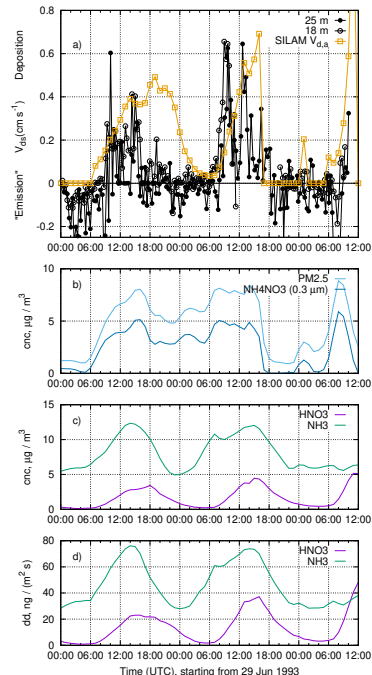
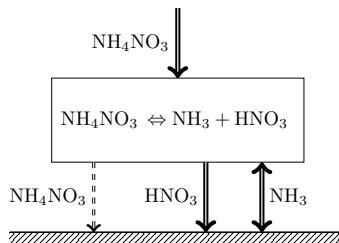
- European air-quality setup
- Primary and secondary inorganic aerosols

NH_4NO_3 flux

Apparent V_d @ obs. height

vs

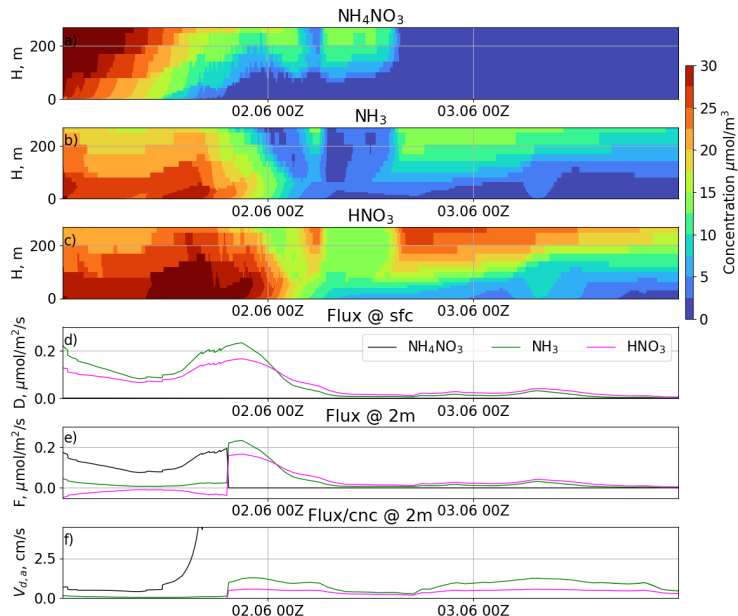
Particle and gas deposition @ sfc



Single-column experiment

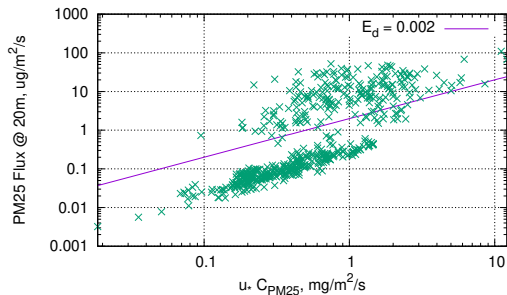
Indirect V_d "measurements"

- Uniform $5 \mu\text{g}/\text{m}^3$ NH_4NO_3 at start
- No wind, no rains
- V_d for aerosol $\sim 0.1 \text{ mm/s}$
- Observe
 - Concentrations
 - Fluxes at 2 m height
 - Depositions (surface)



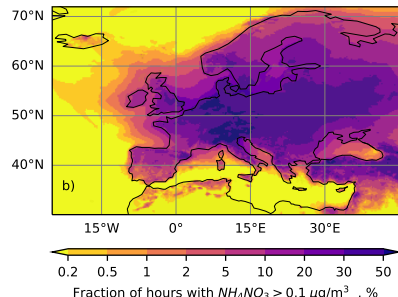
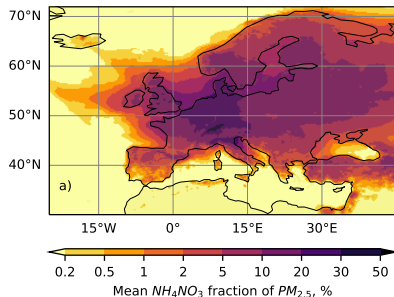
Flux proportional to concentration?

- Scatter of flux vs concentration
- For the simulated case above
- Clusters – present/absent NH_4NO_3
- Never plotted for field studies
- Line – Sulphate ' V_d ' from Weseley (1985)



Abundance of NH_4NO_3

- SILAM CAMS-REG, 2024
- Mean fraction of PM
- Fraction of hours when NH_4NO_3 present



Concluding remarks

- V_d implies that flux is proportional to concentration:

$$\text{Flux} = V_d(\text{particle, flow, surface}) \cdot \text{Concentration}$$

- V_d does not apply to NH_4NO_3 particles
- Same true for many other aerosols
- NH_4NO_3 alone sufficient to break V_d for ambient particles
- SILAM with KS2012 *can* reproduce observed fluxes and apparent V_d
- Accumulation mode deposits at ~ 0.1 mm/s (below detection limit?)
- "Ground-truth"-based models overstate accumulation-mode V_d over forests
- KS2102 scheme ([Kouznetsov & Sofiev, 2012, JGR](#))
 - One deposition-specific parameter of surface for particles
 - Has been used in SILAM since 2011
 - LUC-agnostic